



## Recovery targets and timescales for Lough Neagh and other lakes

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1 **Recovery targets for Lough Neagh and other lakes**

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27 **Abstract**

28

29 The framework, model and methods of Nürnberg were applied and evaluated in Lough  
30 Neagh and 19 other lakes in order to establish inflow phosphorus concentrations that support  
31 target lake values. Supporting concentrations, in the absence of an internal load, were  
32 derived and the effect of uncertainty in the model retention coefficient was relatively small,  
33  $\pm 11-20\%$  in Lough Neagh and an average ( $n=17$ ) of  $\pm 9.7\%$  in the other lakes. There was  
34 further support for the model and methods from an independent estimate of the net internal  
35 load in Lough Neagh (13 % difference) and from another model in the other lakes  
36 (Supporting concentrations, which should be lower, were by an average of  $11 \text{ mg P m}^{-3}$ ). In  
37 the framework, steady state with the phosphorus load is assumed, but, based on a generic lake  
38 model, is not likely if the hydraulic residence time  $> 0.5-0.8$  yr and should lead to a decrease in  
39 phosphorus retention, which was found during three periods in Lough Neagh. Based on a  
40 compilation of internal load recovery times from 23 lakes in the literature, it could take  
41 between 8 and 20 years for lakes with an internal load to achieve their targets.

42

43

44 Key words

45

46 Lakes

47 Phosphorus

48 Targets

49 Good/Moderate boundary

50 Timescales

51 Internal load

## 52 1 Introduction

53

54 Lake nutrient models, especially for phosphorus, have been a major and sustained research  
55 topic in limnology. A lake phosphorus model was an important element in the structure that  
56 linked nutrient sources in the catchment to the biological quality of the lake proposed early in  
57 the field by Dillon & Rigler (1975), a framework that is still valid today. Useful water  
58 quality models also allow managers to assess how a lake should respond to actions in the  
59 catchment and help stakeholders recognize this relationship (Manno et al. 2008).

60

61 While research on this topic has generally focussed on model concepts, complexity and  
62 accuracy (Prairie 1989; Arhonditsis & Brett 2004; Brett & Benjamin 2008; Khorasani & Zhu  
63 2021), the main practical application has been to estimate the nutrient load or concentration,  
64 usually of phosphorus, in the inflow(s) that supports a target lake concentration and  
65 biological condition. Directed management of catchment activities to reduce the export of  
66 phosphorus, nitrogen or both should, ideally, lead to a high quality lake.

67

68 Within the paradigm of the condensed box (Continuously Stirred Tank Reactor) lake model,  
69 which has a large literature, the publications of Prairie (1989), Brett & Benjamin (2008) and  
70 Khorasani & Zhu (2021) cite most of the models. An important issue that needs to be  
71 considered when applying one to a lake that has a high phosphorus concentration, where the  
72 target is to reduce it, is the internal phosphorus load. Generally, this load is high in such  
73 lakes but should decrease as the external load is reduced, albeit slowly (Jeppesen et al. 2005).  
74 Where the internal load is not explicitly incorporated into the model, its effect is assumed to  
75 contribute via the range of lakes used to calibrate the model (Nürnberg 2009). Lakes with a  
76 high phosphorus concentration often have considerable internal loads and *vice versa*, so the

77 varying contribution of internal load to the lake concentration is incorporated through the  
78 range of lakes used to calibrate the model.

79

80 Nürnberg (2009) noted the uncertainty with the contribution of internal load when calibrating  
81 a lake phosphorus model, and the lack of control of it, and went on to formally separate the  
82 contributions of external and internal phosphorus load to the lake concentration (Nürnberg  
83 1984; 1998; 2009). That model predicts the steady-state concentration from the external and  
84 internal loads and she suggested three methods to estimate the internal load, depending on the  
85 type of results available and whether the lake thermally stratifies. This approach can also be  
86 used to derive inflow phosphorus concentrations that support a target lake concentration, and  
87 our aim was to apply and evaluate the framework, model and methods in Lough Neagh and  
88 19 other smaller lakes. The model, as with others used in the field, assumes steady-state and  
89 so, in addition, we assessed the conditions under which a typical CSTR model is at steady  
90 state. Knowing the limitations of a model helps when using it for lake management purposes.

91

92 As the model is a static one, the time it takes to reach steady-state with the inflow(s) depends  
93 only on the lake hydraulic residence time. However, if the lake has an internal load, then,  
94 while the model explicitly incorporates the internal load, how long it would take for that load  
95 to reduce needs to be derived separately. The time for reduction of the internal load has been  
96 estimated using lake models (e.g. Katsev et al. 2006), observations on lakes (e.g. Jeppesen et  
97 al. 2005) and through a model of the diagenesis and burial of labile phosphorus in sediment  
98 (e.g. Rippey et al. 2021). There is some consistency in these estimates, with an overall  
99 indication of one to three decades, but to provide some more direct evidence, we collated  
100 results from lakes that had their external loads reduced and for which sufficient monitoring  
101 results were available.

102

103 Therefore, the aim of the research was to apply and evaluate the framework, model and  
104 methods of Nürnberg (1984; 1998; 2009) in order to derive an inflow phosphorus  
105 concentration that supports a target lake concentration in Lough Neagh and 19 other smaller  
106 lakes, including establishing the uncertainties in estimating the lake concentration from the  
107 external and internal loads. We also estimated the time it should take to reach the target lake  
108 concentrations. There were two supporting objectives: to identify the conditions under which  
109 a typical CSTR model is at steady state and collate results provide direct evidence of the time  
110 it takes for the internal phosphorus load to reduce.

111

## 112 2. Material and methods

113

114 The symbols and definitions used here are given in Table 1. With  $R_{pred}$ , we use the original  
115 definition, net retention in oxic lakes that have no or a naturally low net internal load  
116 (Nürnberg 1998), before the change to gross sedimentation only (Nürnberg 2009).

117 Table 1. Symbols and definitions used.

Property	Definition
$z$ , m	Mean depth
$\tau_w$ , yr	Annual mean hydraulic residence time. $\rho$ , $\text{yr}^{-1}$ , the hydraulic flushing coefficient is $\tau_w^{-1}$ .
$q_s$ , $\text{m yr}^{-1}$	Annual hydraulic load, $q_s = z/\tau_w$
$C_{ext}$ , $\text{mg P m}^{-3}$	Annual volume-weighted average total phosphorus concentration in the inflow(s). $C_{ext} = L_{ext}/q_s$ , where $L_{ext}$ , $\text{g P m}^{-2} \text{ yr}^{-1}$ is the annual average areal external total phosphorus load.
Obs C, $\text{mg P m}^{-3}$	Annual average or monthly average total phosphorus concentration in the lake. Predicted C is the concentration predicted by a lake phosphorus model.
$R_{meas}$ , dimensionless	Implicitly the total phosphorus retention coefficient for one year, is (Input total phosphorus load-Output total phosphorus load)/Input total phosphorus load or, equivalently, Retained load/Input load

Rpred, dimensionless	Implicitly the total phosphorus retention coefficient for one year, predicted by the R7 model of Nürnberg (1984)
Cint, mg P m <sup>-3</sup>	The internal total phosphorus load expressed as a concentration, implicitly in one year. Cint = Lint/qs, where Lint, g P m <sup>-2</sup> yr <sup>-1</sup> , is the annual average areal internal total phosphorus load.

---

118

119 2.1. *Lough Neagh and its phosphorus budget*

120

121 Lough Neagh, a site in the UK's Environmental Change network, is a large (383 km<sup>2</sup>)  
 122 relatively shallow (z 8.9 m) lake, which only rarely stratifies, but has a long history of  
 123 eutrophication (Wood & Smith 1993), up to the 1960s as a result of an increasing external  
 124 phosphorus load (Foy et al. 2003), since then by phosphorus and nitrogen (Bunting et al.  
 125 2007). Since the mid-1990s, a reduction in external nitrogen load and lake concentration  
 126 have resulted in a decrease in the chlorophyll *a* concentration (McElarney et al. 2021; Elliott  
 127 et al. 2016).

128

129 The six main river catchments (Upper Bann, Blackwater, Ballinderry, Moyola, Main and  
 130 Sixmilewater) and two smaller catchments (Crumlin and Glenavy) of Lough Neagh have  
 131 been gauged and monitored since the 1970s, along with the outflow (Lower Bann).  
 132 Discharge is measured continuously (15 min internals) and water samples taken weekly from  
 133 the eight inflows and outflow. The annual mean Input and Output total phosphorus loads are  
 134 calculated using Method 5 in Johnes (2007), i.e. sum of total loads over sampling days  
 135 multiplied by the annual mean discharge as a proportion of the total discharge over the  
 136 sampling days. As the monitored catchments comprise 86 % of the whole catchment, the  
 137 whole catchment load was estimated as load from the monitored catchments divided by 0.86.  
 138 If more than 4 % of flow results for a catchment were missing, the load was not estimated.  
 139 There were no outflow results for 2007 and no results for the Upper Bann in 1995,

140 Sixmilewater in 2010 and Moyola in 1983 and 2010-11, so no catchment phosphorus loads  
141 were estimated for 1983, 1995, 2007 and 2010-2011.

142

143 An integrated water sample was taken at a mid-lake station in 10 m depth, weekly up to 1992,  
144 fortnightly thereafter, and the mean total phosphorus concentration calculated; there were no  
145 lake results for 2009. The Limit of Quantitation for the total phosphorus determination is 10  
146 mg P m<sup>-3</sup>.

147

148 R<sub>meas</sub> was calculated as (Input-Output)/Input and Annual Obs C as the arithmetic mean of  
149 typically 30.3 results in each year, minimum 15 and maximum 52.

150

## 151 2.2. *Other lakes*

152

153 Samples were retrieved at least monthly from the shores of 19 other lakes from 1999, 2004 or  
154 2008 onwards. The Limit of Quantitation of the total phosphorus determination is 10 mg P  
155 m<sup>-3</sup>. The theoretical  $\tau_w$  values were estimated using the measured catchment area, lake area  
156 and mean depth, along with rainfall and potential evapotranspiration estimates made by  
157 interpolation using long-term (1971-2000) data sets of annual mean rainfall and potential  
158 evapotranspiration for sites throughout Northern Ireland.

159

## 160 2.3. *Lake phosphorus models*

161

162 The main use of the models was to estimate the inflow phosphorus concentration (C<sub>ext</sub>) that  
163 supports a target lake concentration (Section 2.6) and the lake phosphorus framework, model  
164 and methods of Nürnberg (1984; 1998; 2009) was used. While there is choice of whether to



165 use the gross internal load or the net, after any sedimentation, we used the net for Lough  
166 Neagh as it could be estimated using the phosphorus mass balance budget (Net Cint\_2)  
167 method (Section 2.5) and, therefore, the model is (Nürnberg 2009):

168

$$169 \text{ Predicted } C = C_{\text{ext}}(1 - R_{\text{pred}}) + \text{Net Cint}_{2} \text{ where } R_{\text{pred}} = 15 / (18 + q_s) \text{ and Net Cint}_{2} = \\ 170 C_{\text{ext}}(R_{\text{pred}} - R_{\text{meas}}) \qquad \qquad \qquad \text{Eqn. 1}$$

171

172 We assumed that a lake of high quality, at or better than the Good/Moderate (Quevauviller et  
173 al. 2008) boundary (Section 2.6), would have little or no net internal load, so Net Cint\_2 in  
174 Eqn. 1 would be zero. Rpred, estimated using the R7 model of Nürnberg (1984), is therefore  
175 key in Eqn.1. That statistical model was developed and calibrated using oxie lakes (n=54)  
176 and so there is some uncertainty in the predicted value for a specific lake. Based on Fig. 1A  
177 in Nürnberg (1984), it applies over a qs range up to 80 m yr<sup>-1</sup> with an uncertainty of ±0.05 up  
178 to a qs value of 20 m yr<sup>-1</sup>, ±0.1 at higher values. An uncertainty in Rpred of ±0.05 was used  
179 when applying Eqn. 1 and when calculating Net Cint\_2 (Section 2.5).

180

181 Even incorporating this uncertainty into the model predictions, some additional evidence to  
182 support the Rpred value or the accuracy of the lake concentrations based on Eqn. 1 with little  
183 or no internal load is desirable (Nürnberg 2009). For this, we used the predictions of another  
184 lake phosphorus model, Model 28 of Khorasani & Zhu (2021). It was selected after  
185 comparing the behaviour of three general lake models and R7 Nürnberg (1984) (Section 1  
186 Supplementary materials) and establishing its accuracy in 14 Irish lakes (Section 2  
187 Supplementary materials). Model 28 of Khorasani & Zhu and Eqn. 9 of Prairie performed  
188 best in these 14 lakes which have some and variable internal load. The mean and median  
189 errors (Observed-Predicted) are -6.2 and -2.3 mg P m<sup>-3</sup> for Prairie and -2.3 and -0.8 mg P m<sup>-3</sup>

190 for Khorasani & Zhu Model 28. The error is larger in some lakes but both models are most  
191 accurate at Predicted  $C < 50 \text{ mg P m}^{-3}$ .

192

193 Based on this assessment, we used Model 28 of Khorasani & Zhu (2021) to provide some  
194 support for the  $C_{ext}$  value derived using Eqn. 1 in the other 19 lakes. As the data set ( $n=738$ )  
195 used to calibrate this model contains lakes with unknown and variable internal loads, the  
196 supporting  $C_{ext}$  concentrations should be lower than those from Eqn. 1 with a Net  $C_{int\_2}$  of  
197 zero, which applies to lakes with no or naturally low internal phosphorus loads. Model 28 of  
198 Khursani & Zhu (2021);

199

$$200 \text{ Predicted } C = CF \times \frac{-1 + \left(1 + 4 \times 0.095 \times t_w^{0.489} \times C_{ext}^{(-0.333+1)} \times Z^{0.288}\right)^{0.5}}{2 \times t_w^{0.489} \times C_{ext}^{-0.333} \times Z^{0.288}} \quad \text{Eqn. 2}$$

201

202 Where CF, the Correction Factor (Sprugel 1983), is 1.159.

203

#### 204 2.4. *Steady-state of lake phosphorus models*

205

206 We investigated the conditions under which a lake is in steady state with its inflow  
207 phosphorus and internal phosphorus cycling, mainly the internal load. In a CSTR model, loss  
208 through the outflow is described by  $\rho$  or  $\tau_w$  and permanent loss to the sediment by the  
209 sedimentation coefficient ( $\sigma$ ) (Chapra 1997, Lecture 3). A time to 90 % of the steady-state  
210 concentration ( $t_{90}$ ) was used, as most of the change has occurred by then, with progressively  
211 slower change to 95 or 99 % of the steady-state value thereafter.  $t_{90}$  for loss only through the  
212 outflow is  $2.303\tau_w$ . As the residence time for sedimentation ( $t_s$ ) is  $\sigma^{-1}$ ,  $t_{90}$  for loss only  
213 through sedimentation is  $2.303\tau_s$ .

214

215 The  $t_{90}$  times, for  $\tau_w$  values of 0.5, 1.0 and 1.5 yr and  $\sigma$  values between 0.3 and 1.0  $\text{yr}^{-1}$  to  
216 represent smaller eutrophic lakes, were first calculated separately and then combined. The  $\sigma$   
217 range was selected based on the following. For populations of lakes, Khorasani & Zhu  
218 (2021) derived a value of 0.786  $\text{yr}^{-1}$ , Jones & Bachman (1976) 0.65 and Brett & Benjamin  
219 (2008) 0.45. In a hierarchical framework, Cheng et al. (2010) estimated values for all lakes  
220 and for categories of  $z$  and  $\tau_w$ ; 0.45  $\text{yr}^{-1}$  for all lakes, but 0.73 for lakes with  $z \leq 10.3$  m and  
221 0.32 for  $>10.3$  m and 1.01 for lakes with  $\tau_w \leq 2.6$  yr and 0.33 for  $>2.6$  yr. In summary,  $\sigma$   
222 values for most lakes vary between 0.3 and 1.0  $\text{yr}^{-1}$ .

223

224 The overall residence time for 90 % of the combined loss through the outflow and by  
225 sedimentation ( $t_{90, \text{combined}}$ ) was calculated using the addition of residence times (Lerman  
226 1979, pp.5);

227

228  $t_{90, \text{combined}} = 1/(\rho + \sigma)$ , with the  $t_{90} = 2.303 t_{\text{combined}}$ .

229

## 230 2.5. *Internal phosphorus load*

231

232 The annual internal phosphorus load for Lough Neagh was estimated using the phosphorus  
233 mass balance method of Nürnberg (2009; 1998) and expressed as a concentration:  $\text{Net Cint}_2$   
234  $= \text{Cext}(\text{Rpred} - \text{Rmeas})$ . To provide some support for these values and to estimate the internal  
235 load in the other 19 lakes, the *in situ* P increase method of Nürnberg (2009; 1998) was used  
236 and the value expressed as a concentration ( $\text{Cint}_1$ ). This method uses the increase of  
237 concentration during the summer and the load lies between the gross and net value.

238

## 239 2.6. *Target lake total phosphorus concentration*

240

241 The target lake total phosphorus concentrations were derived using the WFD-UKTAG (2016)  
242 methodology and was taken to be at the Good/Moderate boundary. With this approach, the  
243 site-specific reference concentration is estimated using statistical models, from which the  
244 boundary values for five quality classes, High/Good, Good/Moderate, Moderate/Poor and  
245 Poor/Bad, are calculated. The statistical models were derived by Cardoso et al. (2007) using  
246 results from 567 lakes distributed throughout Europe and considered to be at reference  
247 condition (only slightly impacted). The reference concentration is the annual geometric mean  
248 value and the statistical models use altitude, the Morphoedaphic Index (alkalinity and mean  
249 depth) and region. There are two regions and one additional model for humic lakes  
250 (Hazen>30 mg/L Pt) in Northern Ireland.

251

## 252 2.7. *Time to reduce the internal phosphorus load*

253

254 To provide additional information on how long it takes the internal load to reduce to a low or  
255 natural amount, we collated results from the literature from lakes that had their external loads  
256 reduced and had been monitored for at least 5 years. The following results were collated  
257 from tables or digitized from charts:  $\tau_w$ ,  $q_s$ , the time period of monitoring after reduction in  
258 external load and  $C_{ext}$  at the start and end of the monitoring. From this information, an  
259 estimate of the time for the internal load to decrease was made. Errors may have been  
260 introduced during digitizing the charts. In the valuable compilation of Jeppesen et al. (2005),  
261 no results for  $C_{ext}$  were directly presented. However,  $C_{ext}$  was estimated from the Predicted  
262 C values in their Fig. 1(a) by applying the lake model in their Eqn. 1.

263

## 264 3. Results

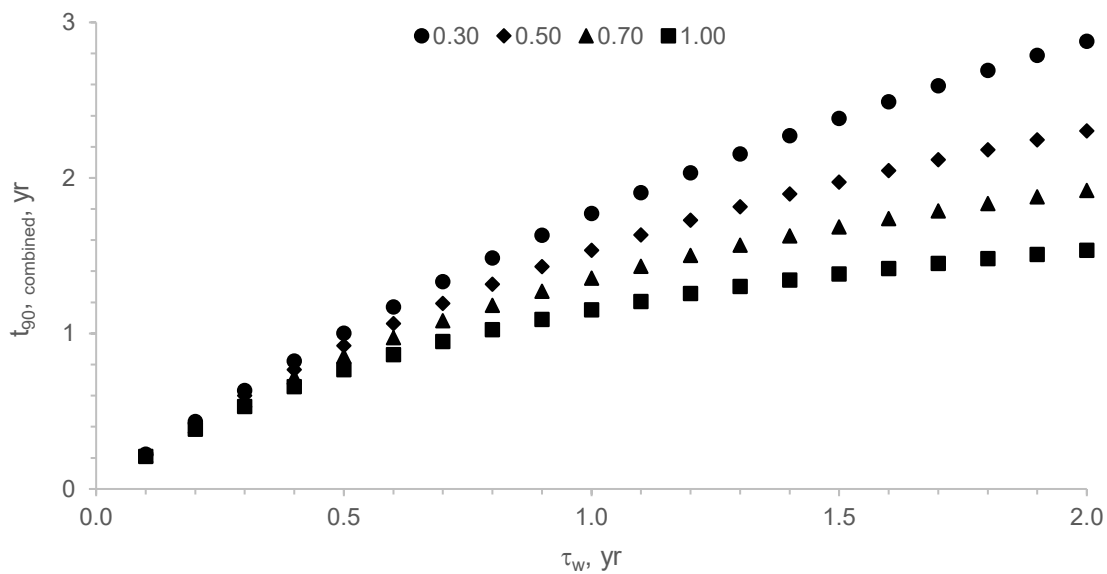
265

266 3.1. Steady state of lake phosphorus models

267

268 As the CSTR model used assumes steady state, the conditions under which this applies were  
269 assessed. The  $t_{90}$  times for loss only through the outflow represented by  $\tau_w$  of 0.5, 1.0 and 1.5  
270 yr are 1.15, 2.30 and 3.45 yr, respectively, loss only through sedimentation represented by  $\sigma$   
271 of 0.3 and 1.0  $\text{yr}^{-1}$  are 7.68 and 2.30 yr, respectively, and the variation of  $t_{90, \text{combined}}$  for the  
272 combined losses is shown in Fig. 1.

273



274

275 Fig. 1. The variation with  $\tau_w$  of the time to 90 % steady state lake water total phosphorus  
276 concentration for combined loss through the outflow and by sedimentation ( $t_{90, \text{combined}}$ ) for four values  
277 of the sedimentation coefficient ( $\sigma$ ) in a Continuously Stirred Tank Reactor.

278

279 For  $t_{90, \text{combined}}$  to be 1 yr or less,  $\tau_w$  needs to be less than 0.5 to 0.8 yr for a lake to be at steady  
280 state with  $\sigma < 1 \text{ yr}^{-1}$ . This should apply to a relatively large change in  $C_{\text{ext}}$  at the start of a

281 year and it is possible this criterion could be relaxed for smaller changes in  $C_{ext}$ . Changes in  
282  $C_{ext}$  later in the year may only produce steady state the following year.

283

284 This is a simple framework with generic  $\tau_w$  and  $\sigma$  values, and phosphorus sedimentation in  
285 smaller rapidly flushed lakes is best described using the retention coefficient (Prairie 1989).

286 However, it suggests that only relatively rapidly flushed lakes would be at steady state with a  
287 changing phosphorus load..

288

289 While there are implications for the estimation of  $\sigma$  in lakes that are not at steady state, the  
290 effect on the more empirically based  $R_{meas}$  is relevant here. In a lake at steady state, loss  
291 through the outflow and to the sediment are at equilibrium with the input, therefore, as the  
292 Retained load is in equilibrium with the Input load,  $R_{meas}$  is accurate (Retained load/Input  
293 load). In a lake not at steady state, if Input is increasing, then loss through the outflow or to  
294 the sediment or both would not be complete and so the Retained load would be smaller than  
295 that at steady state. While  $R_{meas}$  is Retained/Input, a smaller Retained load should lead to a  
296 smaller  $R_{meas}$ . Further, the phosphorus that is in excess of the steady state concentration  
297 would be carried over to the following year, contributing to a smaller  $R_{meas}$  in that year as  
298 well. The same outcome, a depression in  $R_{meas}$ , also applies to a decreasing Input load  
299 when a lake is not at steady state.

300

301 Based on this assessment, a lake is not likely to be in steady state if  $\tau_w > 0.5-0.8$  yr and this  
302 may lead to a decrease in  $R_{meas}$  for one or more years, compared to the steady state value.

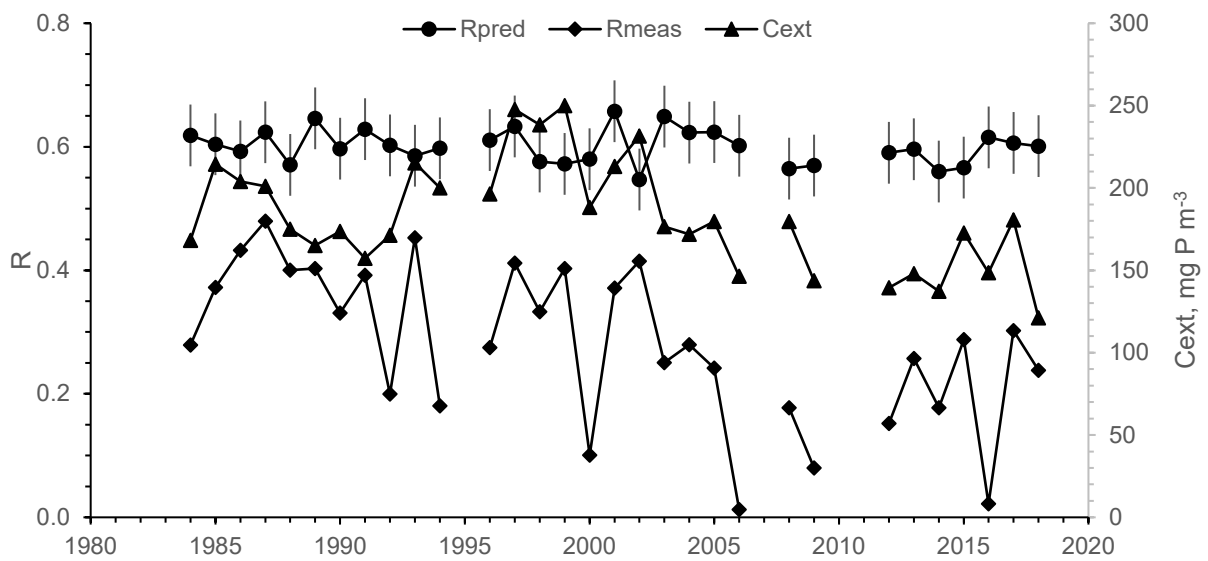
303

304 3.2. Lough Neagh phosphorus budget

305

306 The Lough Neagh phosphorus model (Eqn. 1) uses annual values for  $C_{ext}$ ,  $\tau_w/q_s$ , Input and  
 307 Output total phosphorus load, from which  $R_{pred}$  and  $R_{meas}$  are calculated and then Net  
 308  $C_{int\_2}$  and Predicted C. The variation of  $\tau_w$ , Input load, Output load and Retained load over  
 309 1984-2018 is shown in Section 3 Supplementary materials, with changes in the key  
 310 properties,  $R_{pred}$ ,  $R_{meas}$  and  $C_{ext}$  presented in Fig. 2 and Obs C,  $C_{ext}(1-R_{pred})$  and Net  
 311  $C_{int\_2}$  in Fig. 3.

312



313

314 Fig. 2. The variation of  $R_{pred}$ ,  $R_{meas}$  and  $C_{ext}$  in Lough Neagh, 1984-2018. The error bars  
 315 represent an uncertainty in  $R_{pred}$  of  $\pm 0.05$ .

316

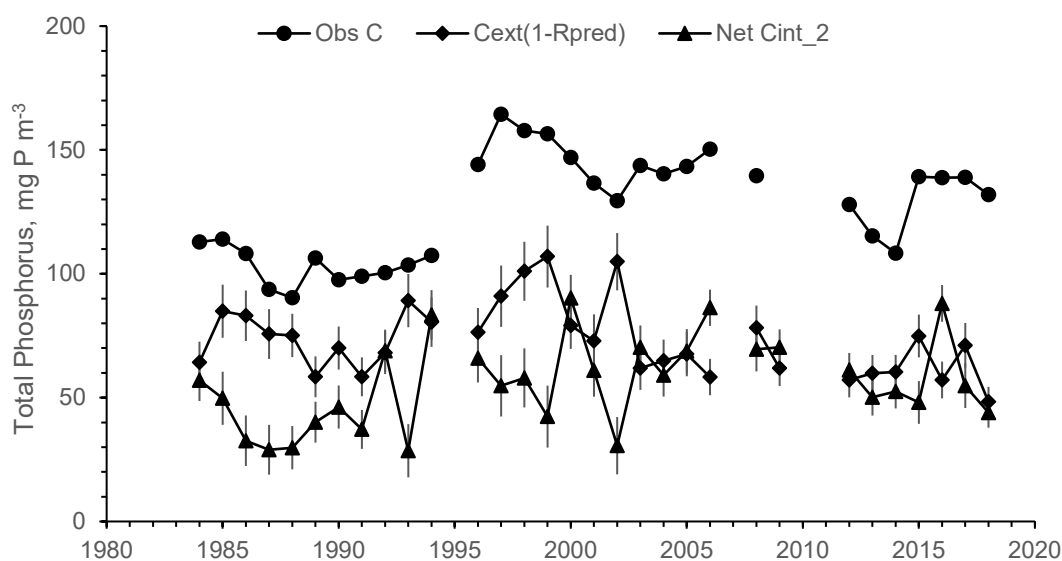
317  $R_{pred}$  varied little, being influenced only by relatively small changes in  $q_s/\tau_w$ , while  $R_{meas}$   
 318 was quite variable. It tended to be smaller from 2003 onwards, with an average of  
 319  $0.357 \pm 0.0298$  between 1984-1994,  $0.330 \pm 0.0427$  between 1995-2002 and  $0.191 \pm 0.0274$   
 320 between 2003-2018 (Table 1 Section 3 Supplementary materials). The reasons for this  
 321 change are investigated in the Discussion.

322

323 In relation to the budget, described fully in Section 3 Supplementary materials, some features  
 324 can be noted. **First**, the variations in Input load are mainly determined by changes in the  
 325 annual inflow volume. **Second**, it follows that the Input and Output loads decrease as  $\tau_w$   
 326 increases, although there is always substantial variability at a given  $\tau_w$  value. **Finally**, even  
 327 with the variability in Input, Output and Retained loads, there were three periods during  
 328 which Input load, Cext, Obs C, **Rpred**, Rmeas, **Net Cint\_2** and  $\tau_w$  were different but relatively  
 329 constant, 1984-1994, 1995-2002 and 2003-2018 (Table 1 Section 3 Supplementary  
 330 materials).

331  
 332 As the model is Predicted C = Cext(1-Rpred) + Net Cint\_2 (Eqn. 1), Fig 3 shows how the  
 333 relative contributions of **the net** external and **net** internal loads to the lake water concentration  
 334 vary. Both contribute, with external load more important up to 1999, generally equal  
 335 afterwards. There **were** rapid year to year variations in both contributions, **particularly**  
 336 **between 1991 and 1994, 1998-2003 and 2014-2017** and especially **with** internal load. **The**  
 337 **variations in internal load were mainly caused by changes in Rmeas (Fig. 2), as Net Cint\_2 =**  
 338 **Cext(Rpred-Rmeas), with Rpred relatively constant and Cext intermediate (Fig. 2).**

339



340



341 Fig. 3. The variation of Obs C, Cext(1-Rpred) and Net Cint\_2 in Lough Neagh, 1984-2018. The error  
342 bars represent the effect of an uncertainty in Rpred of  $\pm 0.05$ .

343

344 The reason for these rapid changes in Net Cint\_2 and Rmeas (1991-1994, 1998-2003, 2014-  
345 2017) is likely to the lake not being at steady state with changing Input phosphorus load.

346 First, the overall average  $\tau_w$  is 1.30 yr (Section 3 Supplementary materials), larger than the  
347 criterion of  $<0.6-0.8$  yr suggested for steady state. Second, a detailed analysis of the changes  
348 in Rmeas and Input phosphorus load (Section 4 Supplementary materials) shows that the  
349 reductions in Rmeas coincide with a rapid change in the Input load eight out of ten times  
350 during these three periods.

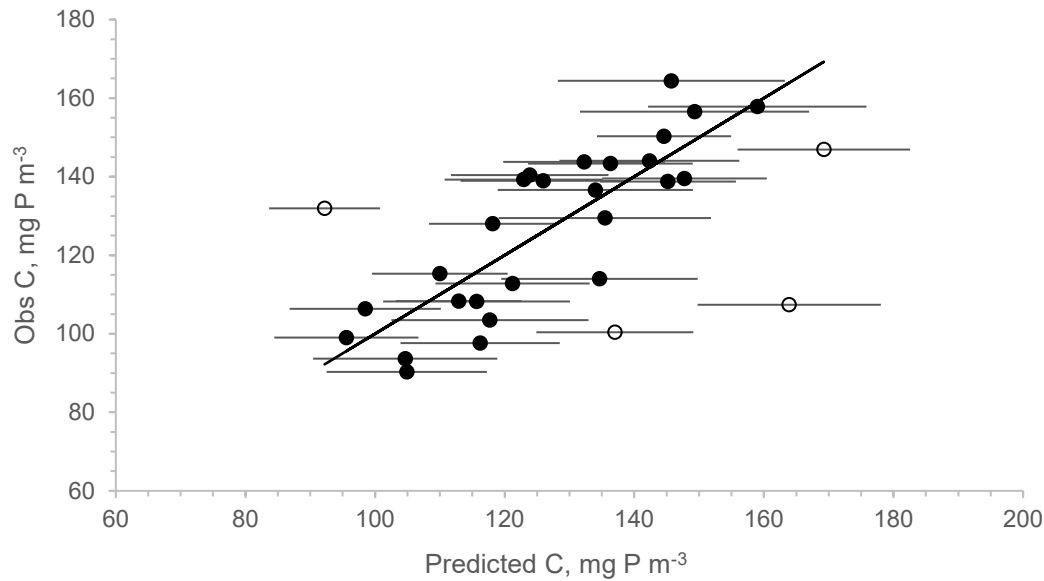
351

352 3.3. Lough Neagh phosphorus model and supporting Cext

353

354 The relationship between Obs C and Predicted C, 1984-2018, is shown in Fig. 4. The  
355 correlation coefficient is 0.616 ( $p < 0.001$ ), with a mean and median error (Obs C - Predicted C)  
356 of -2.3 and 0.28 mg P m<sup>-3</sup>. While there is almost no bias over the 30 years of predictions, the  
357 error is relatively large ( $>20$  mg P m<sup>-3</sup>) in 1992, 1994, 2000 and 2018. The first three years  
358 fall within two of the periods of rapid changes in Rmeas and Net Cint\_2, 1991-1994, 1998-  
359 2003, and the fourth is close to the third period (2014-2017). In 2018, the Predicted C (92.2  
360 mg P m<sup>-3</sup>) and Cext(1-Rpred) (48.3 mg P m<sup>-3</sup>) were the lowest in the monitoring period,  
361 while Obs C was not unusual (Fig. 3), leading to the large error. Omitting these four years,  
362 the correlation coefficient becomes 0.860 ( $p < 0.001$ ) and the mean and median errors 0.21 and  
363 2.2 mg P m<sup>-3</sup>.

364



365

366 Fig.4. Relationship between Annual Obs C and Predicted C in Lough Neagh, 1984-2018, based on  
 367 the lake phosphorus model in Eqn. 1. The error bars are the propagation of the errors for external  
 368 load ( $C_{ext}(1-R_{pred})$ ) and net internal load ( $C_{ext}(R_{pred}-R_{meas})$ ) based on the effect of an uncertainty  
 369 in  $R_{pred}$  of  $\pm 0.05$ . The unity line is shown and the open circles are 1992, 1994, 2000 and 2018.

370

371 The main evidence available to support this model is another estimate of the internal load.

372 The *in situ* P increase method ( $C_{int\_1}$ ) method was applied and the value compared to the

373 results from the phosphorus mass balance method ( $Net\ C_{int\_2}$ ) (Fig. 3). Gibson et al. (2001),

374 in a description and analysis of the internal phosphorus load in Lough Neagh 1974-1997,

375 presented the variation of monthly mean Obs C values for 1995. During the summer, Obs C

376 increased from 70 mg P m<sup>-3</sup> in June to 200 in September, an increase of 130 mg P m<sup>-3</sup>. This

377 *in situ* P increase should lie between the gross and net concentration, but if it is taken to be

378 gross, then  $Net\ C_{int\_1} = (1-R_{pred})Gross\ C_{int\_1}$ . The inflow was missing for one catchment

379 in 1995, so no qs value is available, but the mean of 1994 and 1996 was used for  $R_{pred}$

380 (0.604), giving a  $Net\ C_{int\_1}$  of 51.5 mg P m<sup>-3</sup>. A  $Net\ C_{int\_2}$  value is missing for 1995, and

381 the values fluctuate between 28.6 (1993), 83.4 (1994), 65.9 (1996) and 54.8 (1997), but the

382 mean (58.2 mg P m<sup>-3</sup>) is close to the  $Net\ C_{int\_1}$  value of 51.5, a 13 % difference. This

383 relatively small difference allows us to take the  $C_{int\_1}$  value in Lough Neagh to be an  
384 estimate of the gross internal load. Similar analysis of other years would provide additional  
385 evidence on the nature of the internal load derived by the *in situ* P increase method in Lough  
386 Neagh.

387

388 With this support for the phosphorus model for Lough Neagh, it can be used to derive the  
389  $C_{ext}$  value that supports the (geometric) mean total phosphorus concentration at the  
390 Good/Moderate (G/M) boundary with no or a naturally low net internal phosphorus load.  
391 Lough Neagh is on the boundary between a clear water and a coloured lake, so the G/M  
392 boundary could be 24 or 44 mg P m<sup>-3</sup>, respectively. Based on Eqn. 1, with Net  $C_{int\_2}$  of zero  
393 and the mean  $R_{pred}$  for 2003-2018 of 0.597 (Table 1 Section 3 Supplementary materials),  
394 uncertainty  $\pm 0.05$ , the supporting  $C_{ext}$  concentrations are  $60 \pm 6.6$  ( $\pm 11\%$ ) and  $109 \pm 0.22.1$   
395 ( $\pm 20\%$ ) mg P m<sup>-3</sup>. A lake at the boundary between two classes always needs further  
396 assessment, but Lough Neagh on the boundary between a clear and coloured water lake has a  
397 large effect on the targets for the lake and inflow.

398

399 As the value of  $R_{pred}$  is central to the model (Eqn. 1), we used an uncertainty in  $R_{pred}$  of  
400  $\pm 0.05$  (Fig. 2) to indicate its influence on the estimates of the net external and net internal  
401 loads (Fig. 3). The uncertainties are not large compared to the values of the properties and  
402 the magnitude of the changes. When the errors are combined in the model, as Lough Neagh  
403 has a substantial internal load in its current state, the uncertainty is larger (Fig. 4).

404

405 3.4. The other lakes

406

407 These 19 lakes do not have phosphorus budgets so the monitoring results were used to  
408 describe the long-term change in Annual Obs C and the long-term annual cycle of Monthly  
409 Obs C. The full results are presented in Section 5 Supplementary materials, with two  
410 representative examples shown in Fig. 5 and 6. For each lake, the G/M total phosphorous  
411 boundary was derived, depending on whether it is clear or coloured water, and the Cext value  
412 that supports it was calculated using two models, Eqn. 1, with a Net Cint of zero and an  
413 uncertainty in Rpred of  $\pm 0.05$ , and Eqn. 2.

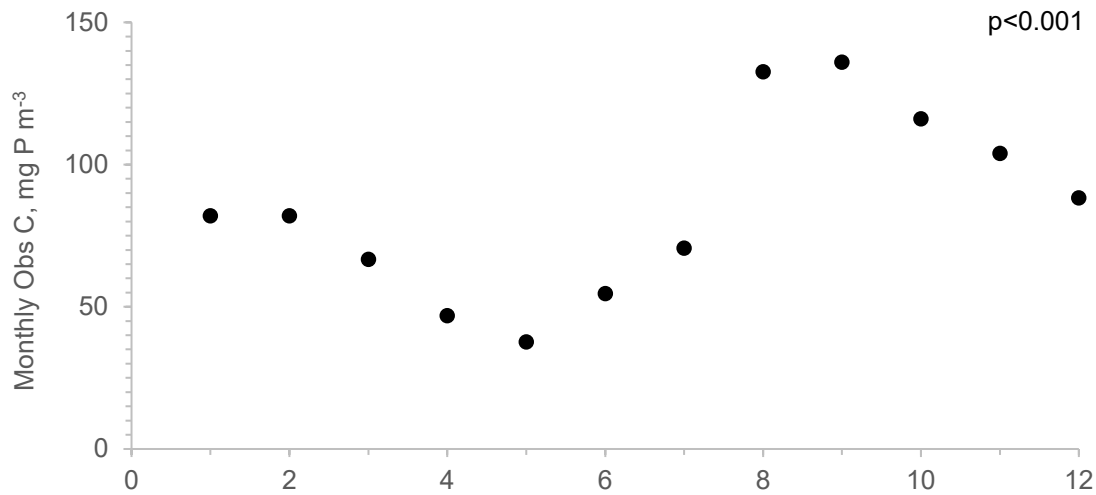
414

415 There was little change in the Annual Obs C in Stoneyford, although a suggestion of an  
416 increase from 2017 (Fig. 5), but the long-term (2004-2018) average concentration (mean of  
417 the Annual Obs C) is  $85.9 \text{ mg P m}^{-3}$ . There is clear evidence of a summer increase in  
418 concentration, due to an internal load, from which a Cint\_1 value of  $95.0 \text{ mgP m}^{-3}$  can be  
419 estimated. It is a coloured water lake with a G/M boundary of  $21 \text{ mg P m}^{-3}$  and the  
420 supporting Cext value is  $42 \pm 4.3 \text{ mg m}^{-3}$  with Eqn 1 and  $29 \text{ mg P m}^{-3}$  with Eqn. 2.

421



422

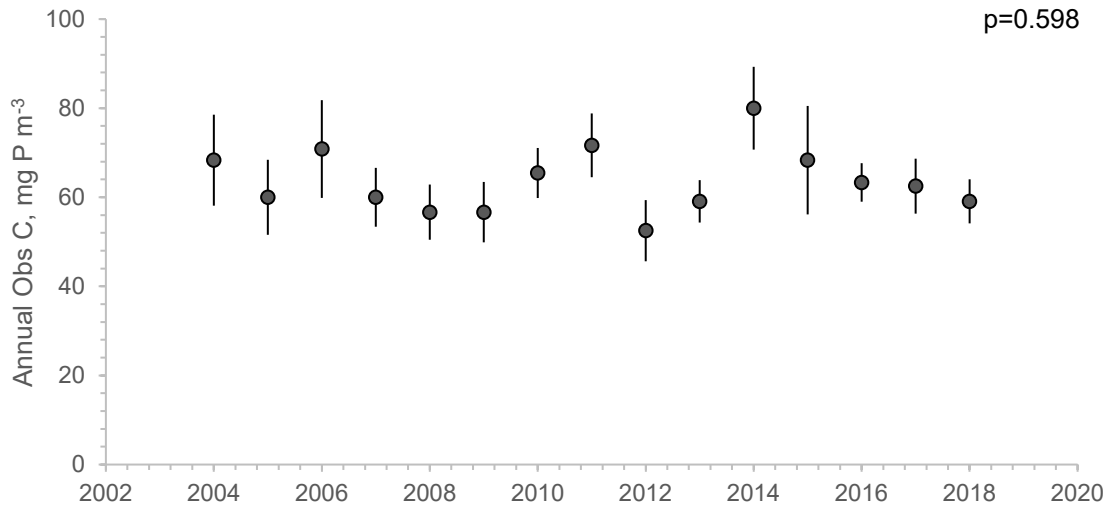


423

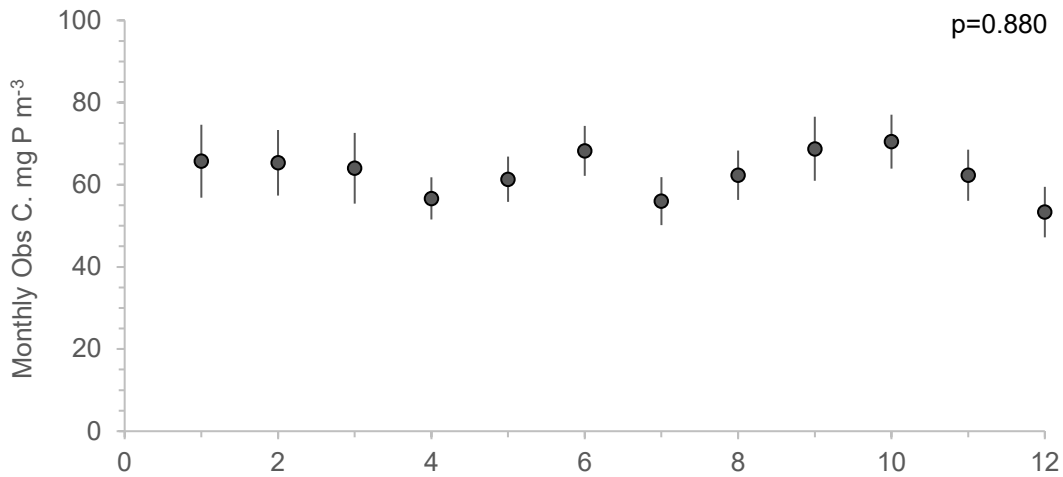
424 Fig. 5. The variation of annual mean total phosphorus concentration (Annual Obs C) in Stoneyford,  
 425 2004-2018, and of the long-term average monthly concentration (Monthly Obs C). The significance  
 426 (p) of a one-way ANOVA is shown.  $\tau_w$  is 0.44 yr and z 5.2 m.

427

428 The was no change in the Annual Obs C in Lough Ross, giving a long-term average  
 429 concentration of 63.6 mg P m<sup>-3</sup>, and there is no evidence for a summer increase in  
 430 concentration (Fig. 6). It is a coloured water lake with a G/M boundary of 37 mg P m<sup>-3</sup> and  
 431 the supporting C<sub>int</sub> is 45±2.7 mg m<sup>-3</sup> with both Eqn. 1 and Eqn.2. Unlike Stoneyford, the  
 432 two models give identical supporting C<sub>ext</sub> values, but this is may be due to the lake's short  $\tau_w$   
 433 time (0.08 yr compared to 0.44 in Stoneyford) increasing the Predicted C in the Khorasani &  
 434 Zhu model (Section 1 Supplementary materials), and consequently producing a larger C<sub>ext</sub>  
 435 value.



436



437

438 Fig. 6. The variation of annual mean total phosphorus concentration (Annual Obs C) in Lough Ross,  
 439 2004-2018, and of the long-term average monthly concentration (Monthly Obs C). The significance  
 440 (p) of a one-way ANOVA is shown.  $\tau_w$  is 0.08 yr and z 5.6 m.

441

442 Table 2 summarize the results for the 19 lakes; note two sites are used for Lower Lough Erne.  
 443 Eleven lakes have no internal load (Cint\_1), eight have, varying from 8.3 mg P m<sup>-3</sup> in Lough  
 444 Island Reavy to 327.5 in Portmore, but typically 20-60 mg P m<sup>-3</sup>. The G/M boundary  
 445 concentration varies from 15 mg P m<sup>-3</sup> in Island Reavy to 65 in Gullion, with a mean of 37.3  
 446 mg P m<sup>-3</sup>, median 36. Five lakes have G/M values greater than 50 mg P m<sup>-3</sup>. They are  
 447 coloured water, shallow ( $z < 2.4$  m) lakes with alkalinities greater than 1.2 meq L<sup>-1</sup>, all of

448 which increase the reference concentration (WFD-UKTAG 2016). Lough Beg is a widening  
449 of the Lower River Bann outlet of Lough Neagh and is omitted from further analysis. There  
450 are two other linked lake systems. Upper Lough MacNeaen (G/M boundary  $37 \text{ mg P m}^{-3}$ )  
451 flows into Lower Lough MacNeaen (G/M boundary  $53 \text{ mg P m}^{-3}$ ) and so achieving the quality  
452 target in the Upper Lough should lead to that in the Lower Lough as well. Similarly, Upper  
453 Lough Erne ( $51 \text{ mg P m}^{-3}$ ) flows into Lower Lough Erne A and B ( $36 \text{ mg P m}^{-3}$ ), but, as the  
454 lower lake is deeper its G/M boundary concentration is lower.

455

456 The Cext values that support the G/M boundary concentrations, in the absence of an internal  
457 load, vary from  $28 \pm 2.1$  to  $80 \pm 10.2 \text{ mg P m}^{-3}$ , except for Clea Lakes ( $148 \pm 21.1$ ), Gullion  
458 ( $336 \pm 93.2$ ) and Portmore ( $134 \pm 14.5$ ). They all have boundary values greater than  $50 \text{ mg P}$   
459  $\text{m}^{-3}$  and higher (0.536-0.807) than the average Rpred value (0.494), both of which increase  
460 the supporting Cext concentration. The effect of the uncertainty in the Rpred value is  
461 relatively small in most of the lakes, only over  $\pm 20 \%$  in two, Gullion ( $\pm 28 \%$ , Rpred 0.807)  
462 and Island Reavy ( $\pm 21 \%$ , 0.748). Both have high Rpred values, which lead to low (1-Rpred)  
463 values, where that the influence of a fixed uncertainty on the Cext is greatest. Ignoring these  
464 two lakes, the mean (n=17) error is 9.7 %, median 9.3, minimum 6.0, maximum 14.

465

466 The only support for the model (Rpred value) in these lakes was the predictions of another  
467 model, Model 28 Khorasani & Zhu (2021). As it contains some and unknown contributions  
468 from internal load, the supporting Cext values should be lower. The difference (Eqn. 1, with  
469 Net Cint<sub>2</sub> of zero, minus Khorasani & Zhu) is over  $50 \text{ mg P m}^{-3}$  in three lakes (Clea lakes  
470 65, Gullion 226, Portmore 64), which have high G/M boundary concentrations. Omitting  
471 these, the mean difference (n=16) is  $12 \text{ mg P m}^{-3}$ , median 14, minimum -9 (Lower Lough

472 Erne), maximum 35 (Fea). There is little basis to determine if these differences are sufficient  
473 to support the Rpred values, but, at least, they are not large for most of the lakes.



474 Table 2. The total phosphorus concentration at the Good/Moderate (G/M) boundary, theoretical hydraulic residence time ( $\tau_w$ ), measured mean depth,  
 475 long-term (1999- or 2004-2018) or recent measured average total phosphorus concentration (Obs C), *in situ* P increase and the supporting inflow  
 476 concentration (Cext) that supports the G/M boundary concentration according to two lake phosphorus models. An uncertainty in R of  $\pm 0.05$  was used with  
 477 the Nurnberg model and is shown in brackets. The timescale of recovery is selected using the criteria in Section 4.3.

Lake	G/M boundary, mg P m <sup>-3</sup>	Theoretical $\tau_w$ , yr	Mean depth, m	Obs C, mg P m <sup>-3</sup>	<i>In situ</i> P increase, mg P m <sup>-3</sup>	Modelled supporting Cext, mg P m <sup>-3</sup>		Timescale, yr
						Nurnberg (1984) R7	Model 28 Khorasani & Zhu (2021)	
Beg	63	Note 1	2.0	136.5	51.4	---	---	Rapid
Cam	18	0.76	5.4	59.6	20.5	45 $\pm$ 5.6	26	8 to 18
Castlehume	28	0.24	3.1	33.3	None	55 $\pm$ 5.	36	0.55
Clea Lakes	53	0.41	2.2	44.5	42.9	148 $\pm$ 21.1	83	8 to 18
Lower Erne		0.33	11.9					0.76
Devenish	36			63.3	None	50 $\pm$ 3.5	59	
Ross Harbour	36			45.8	None	50 $\pm$ 3.5	59	
Upper Erne	51	0.08	2.3	65.1	None	75 $\pm$ 5.6	76	0.18
Fea	32	0.40	2.8	23.6	None	80 $\pm$ 10.2	36	0.92
Gullion	65	1.01	0.6	253.3	60.9	336 $\pm$ 93.2	74	8 to 18
Island Reavy	15	3.82	7.8	30.3	8.3	60 $\pm$ 12.3	23	8.8
	37	0.3	4.4	32.3/23.9	None	68 $\pm$ 6.3	54	0.69
Upper MacNean				Note 2				
	53	0.04	1.4	59.9/31.9	None	74 $\pm$ 5.2	58	0.09
Lower MacNean				Note 2				

Melvin	40	1.05	8.3	17.0	None	95±11.4	83	2.42
Mourne	20	0.80	5.4	63.2	60.4	51±6.5	31	8 to 18
Portmore	62	0.07	0.7	433.0	327.5	134±14.5	70	8 to 18
Ross	37	0.08	5.6	63.6	None	45±2.7	45	0.18
Scolban	35	0.47	7.8	35.8	None	62±5.5	58	1.08
Silent Valley	19	0.26	7.4	24.2	None	28±2.1	25	0.63
Spelga	24	0.41	6.5	26.8	None	43±3.9	35	0.94
Stoneyford	21	0.44	5.2	85.9	95.0	42±4.3	29	8 to 18

478 Note 1. Lough Beg is a widening of the Lower River Bann River exit of Lough Neagh and so its total phosphorus concentration should be identical to Lough  
479 Neagh's.

480 Note 2. Where open lake samples were available, as well as shore samples, the average concentration of shore/open lake are given.

481 Note 3. With a small and uncertain internal load ( $8.3 \text{ mg P m}^{-3}$ ) and a long  $\tau_w$  value (3.82 yr), a  $t_{90}$  of  $2.303\tau_w$  was used.

482 3.5. *Time to reduce the internal phosphorus load*

483

484 The detailed results collated are provided in Section 6 Supplementary materials, with a  
485 summary in Table 3. It illustrates the uncertainty of the estimates, e.g. more than 5-8 years,  
486 over 10 years etc., and there was no relationship between the time and either Start Cext, End  
487 Cext or the change. The estimates are relatively imprecise, as factors such as how quickly  
488 and by how much the external load was reduced, the frequency and length of the monitoring,  
489 length of the initial eutrophication of the lake, size of the internal load, etc. could affect the  
490 time and will vary from lake to lake. However, there is a broad tendency to three values,  
491 approximately 8 (<6 to 10) yr, 14 yr (12 to 16) or 20 yr (Fig. 7); the average (n=21) is 12.2 yr,  
492 7.6 yr for those lakes (n=11) with time <11 yr and 17.2 for those lakes (n=10) with a time 12  
493 yr or more.

494

495

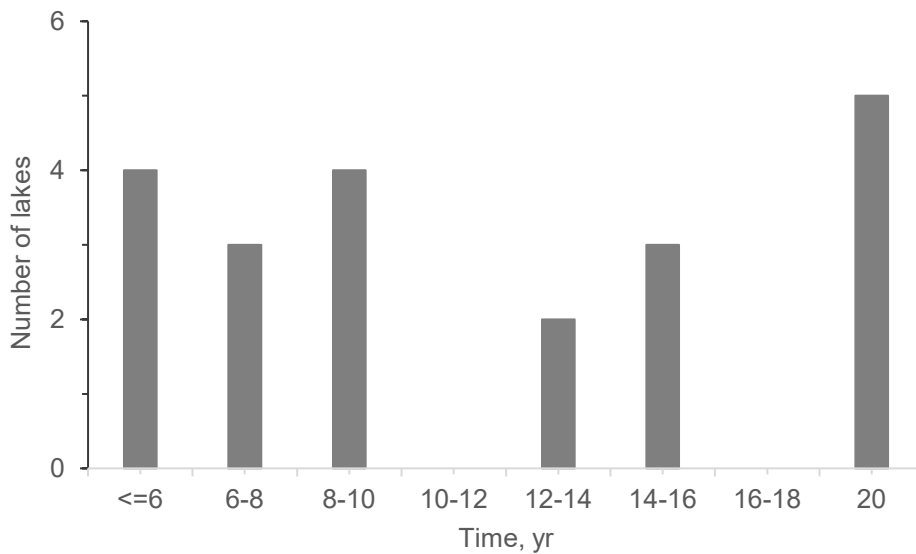
496 Table 3. Summary of the findings of the observed decrease in internal phosphorus load in lakes. The hydraulic residence time ( $\tau_w$ ), hydraulic load (qs), time  
 497 period of results, annual average inflow concentration (Cext) at the start and end of the period and the time for the internal load to decrease are given. If a  
 498 range of  $\tau_w$  and qs values are available, the ranges are given.

Lake	$\tau_w$ , yr	qs, m yr <sup>-1</sup>	Time period	Cext at start and end, mg P m <sup>-3</sup>	Time for internal load to decrease	Reference
Shagawa Lake	0.52-1.2	4.5-10.9	1971-76	72 & 21	5-6 years	1
Oxundasjon	0.15	22.5	1979-75	844 & 93	More than 5-8 years	2
Edssjon	0.18	16.8	1970-75	655 & 131	More than 5-8 years	2
Norrviken	1.1	4.7	1969-78	321 & 74	5-8 years	2
Lake Washington	1.7-3.1	10.5-19.4	1962-78	162-48	7-13 years	3
Mondsee	1.7	21.2	1978-89	51 & 22	Small internal load reduced over 9 years	4
Lake Pyhajarvi	3.0	1.8	1980-92	78 & 78	No internal load	5
Grundsomagle	0.08	15.0	10 years	3079 & 180	15 or more years	6
Albufera	0.1	12.0	10 years	1843 & 421	15 or more years	6
Eemmeer	0.05	34.0	15 years	1224 & 281	15 to 20 years	6
Gooimeer	0.16	21.9	15 years	1008 & 336	Over 10 years	6
Arreso	2.2	1.4	10 years	1167 & 236	13-14 years	6
Vesterborg	0.05	26	5 years	245 & 116	More than 5 years	6
Barton Broad	0.045	33.3	15 years	364 & 97	15 years	6
Sogard	0.05	32.0	10 years	318 & 196	Internal load increased over 10 years	6
Okeechobee	2.7	1.0	20 years	581 & 238	Internal load increased from year 10-20	6
Ornso	0.05	80.0	10 years	208 & 110	Small internal load recovering in 20 years	6
Vortsjarv	1.0	2.8	10 years	240 & 48	Internal load developed in last 5 years	6
Galten	0.05	68.0	20 years	67 & 29	Internal load maintained for 20 years	6

Bryrup	0.15	30.7	10 years	132 & 101	Internal load only at the start, so rapid (5 yr) reduction	6
Loch Leven	0.42	9.3	1975-2005	135 & 154	Internal load began after 1985, maintained and reduced up to 2008. Responds rapidly (few years). Little reduction in Cext	7
Muggelsee	0.12-0.23	21.2-42.6	1979-2016	149 & 67	Internal load between 1982 and 2015 changes rapidly (few years) to changes in lake concentration	8
Lake Apopka	2.5	0.64	1989-2002	1563 & 344	Increasing, small internal load from 1992 to 2002, so more than 10 years	9

499 Baddeggersee ( $\tau_w$  4.2 yr), Sempachersee (16), Hallwillersee (3.9) and Pfaffikersee (2.2) are all calcareous lakes and had reductions in Cext of 40, 30, 66 and  
500 75 %, while Obs C decreased by 95, 80, 90 and 90 %, respectively (Muller et al 2014). The models used the sedimentation coefficient and it was variations in  
501  $\sigma$  that mainly determined Obs C. As Obs C changed over a 30 year period, it is difficult to establish how quickly s changed. However, the rate of change of s  
502 with Obs C increases as Obs C falls below 60-80 mg L<sup>-1</sup>. 1 Larsen et al. (1979). 2 Ahlgren (1980). 3 Edmondson & Lehman (1981). 4 Dokulil & Jagsch (1992).  
503 5 Ekholm et al. (1997). 6 Jeppesen et al. (2005). 7 Spears et al. (2012), May et al. (2012) and Bailey-Watts & Kirika (1999). 8 Shatwell & Kohler (2019) and  
504 Kohler et al. (2005). 9 Coveney et al. (2005)

505



506

507 Fig. 7. Frequency distribution of the observed time for the internal load to reduce in 21 lakes (Table  
508 3).

#### 509 4. Discussion

510

511 The main aim of the research was to apply and evaluate the framework, model and methods  
512 of Nürnberg in order to derive an inflow phosphorus concentrations that support target lake  
513 concentrations, using Lough Neagh and 19 other smaller lake as examples. In the evaluation,  
514 the uncertainties in estimating the lake concentration from the external and internal loads  
515 were estimated and we also identified the conditions under which a typical CSTR model is at  
516 steady state and collated results to provide direct evidence of the time it takes for the internal  
517 phosphorus load to reduce.

518

##### 519 4.1. Nürnberg's framework, model and methods

520

521 The framework to separate the contributions of the external and internal loads to the lake  
522 phosphorus concentration was clear to understand, the CSTR lake model is well established

523 and the methods used to estimate the internal load were easy to apply. Generally, steady state  
524 conditions are assumed, but we identified using a generic CSTR lake model that it is not  
525 likely if  $\tau_w > 0.5-0.8$  yr and should lead to a decrease in  $R_{meas}$  for one or more years,  
526 compared to the steady state value (Section 3.1). We found this behaviour in Lough Neagh,  
527 where there were rapid changes in  $R_{meas}$  during 1991-1994, 1998-2003 and 2014-2017  
528 when the lake was not at steady state with a changing Input phosphorus load (Section 3.2).  
529 Averaging results is recommended when applying these models (Nürnberg 2009).

530

531 The value of  $R_{pred}$  is key in the lake model (Eqn. 1) and so we included a fixed uncertainty  
532 (Fig. 2) in estimating the net external and net internal loads (Fig. 3). The uncertainty is small  
533 compared to the load values and their variation, but a little larger when the external and  
534 internal loads are combined in the model (Fig. 4). However, as there are also uncertainties in  
535 estimating  $C_{ext}$  and  $R_{meas}$  when calculating the loads, its influence may not be the most  
536 important. We also provided some additional support for the model in Lough Neagh (4.1.1)  
537 and the other lakes (4.1.2).

538

#### 539 4.1.1. Lough Neagh

540

541 The Lough Neagh phosphorus model reproduced the variation of annual mean Obs C over  
542 1984-2018 period well, with mean and median errors of -2.3 and 0.28 mg P m<sup>-3</sup>, respectively  
543 (Section 3.3). While this is not an independent assessment of the model, it does provides  
544 some evidence that it can be used to derive a  $C_{ext}$  that supports a target lake concentration.

545

546 There was evidence to support the lake model, in the agreement between two estimates of the  
547 internal load, 51.5 and 58.2 mg P m<sup>-3</sup> (Section 3.3). The difference, 13 %, is relatively small

548 and provides support for the lake model. Foy et al. (2003) reconstructed the phosphorus  
549 budget for Lough Neagh, 1840-2000, using the diatom and chironomid record in sediment  
550 cores and a contemporary budget, identified that sediment release began in the 1960s and  
551 estimated it to be at least  $36 \text{ mg P m}^{-3}$ . If this is taken to be net release and that it would  
552 increase as eutrophication of the lake developed, then this is some further support for the Net  
553 Cint\_2 values in Fig. 3.

554

#### 555 4.1.2. *The other lakes*

556

557 The water column monitoring results were used to establish long-term trends and if there was  
558 an internal load, which was estimated using the *in situ* P increase method (Cint\_1). Cext  
559 values that support the target concentrations were able to be derived (Table 2).

560

561 The only support for the model was to use another lake model (Model 28 of Khorasani & Zhu  
562 2021) to derive supporting Cext values; these should be lower, as that model includes  
563 variable contributions from internal load. For most of the lakes (16 out of 19), the values  
564 were an average of  $12 \text{ mg P m}^{-3}$  lower (Section 3.4), which lends some general support to the  
565 Rpred values used in Eqn. 1 for these lakes.

566

#### 567 4.2. *Time to reduce the internal load and reach steady-state concentration*

568

569 The model is a steady state one and only indicates the time it would take the lake adjust to its  
570 external load based on the hydraulic residence time. If there is an internal load, which is  
571 common in lakes needing remediation, the time it takes for it to reduce is also involved.  
572 Based on observational evidence, there is a general tendency for it to be either approximately



573 8 (<6 to 10) yr, 14 yr (12 to 16) or 20 yr (Fig. 7). These values are similar to the one to three  
574 decades summarized by Rippey et al. (2021) from a compilation of evidence from models  
575 and observations. A better basis to choose a time for internal load reduction in a lake is  
576 desirable, but, in the absence of a direct estimate of 8 to 20 yr can be suggested. If there is no  
577 or little internal load, then the time to the new steady-state concentration would be  
578 determined by  $\tau_w$ , with  $t_{90} = 2.303\tau_w$ . These two rules were used to select the time to steady  
579 state in the 19 lakes (Table 2).

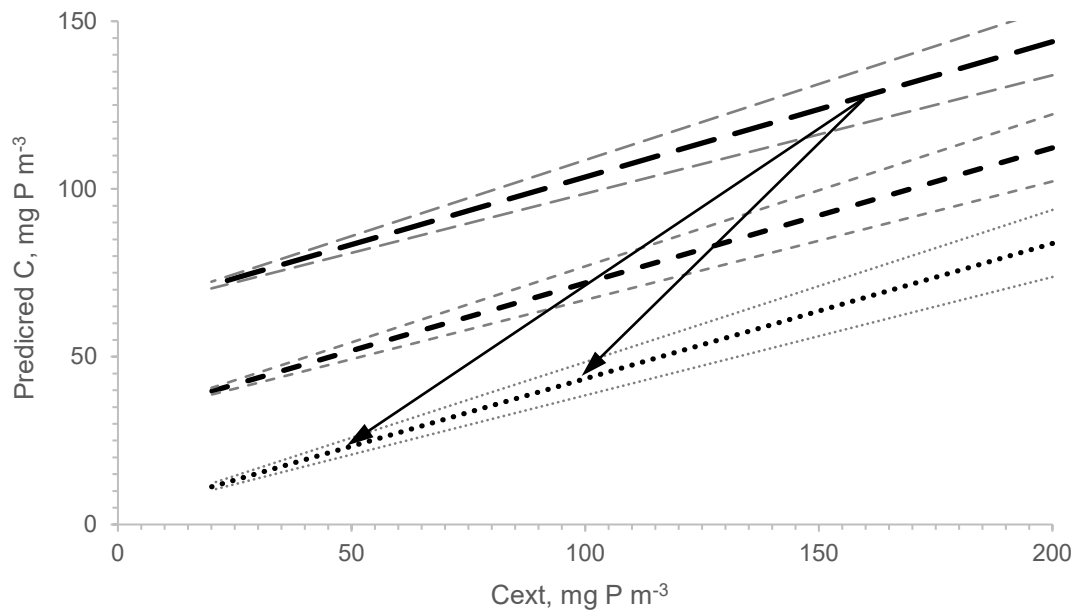
580

#### 581 4.2.1. Lough Neagh

582

583 The  $t_{90}$  value for the internal phosphorus load to reduce in Lough Neagh has been estimated  
584 by applying the Lewis/Penn model to five sediment cores (Rippey et al. 2021), varying from  
585 15 to 23 yr, mean 20.6. As the mean  $\tau_w$  is 1.30 yr, the time to reach the target lake  
586 concentration of 24 or 44 mg P m<sup>-3</sup> should be determined mainly by the reduction in internal  
587 load, so around 21 years. Fig. 8 shows the modelled change in lake phosphorus concentration  
588 to reductions in the inflow concentration for three values of net internal load, as the current  
589 value reduces to a half and one tenth over 21 years. As advised (Nürnberg 2009), average  
590 values were used to apply the model (Eqn. 1); the 2003-2018 mean  $R_{pred}$  is 0.597, Net  
591  $C_{int\_2}$  63.3 mg P m<sup>-3</sup> and  $C_{ext}$  159 mg P m<sup>-3</sup> (Table 1 Section 3 Supplementary materials).  
592 An uncertainty in  $R_{pred}$  of  $\pm 0.05$  was applied.

593



594

595 Fig. 8. The variation of steady-state lake total phosphorus concentration predicted by Eqn. 1  
 596 (Predicted C) with inflow concentration (Cext) for three values of net internal load, full Net Cint\_2  
 597 (2003-2018 average of 63.3 mg P m<sup>-3</sup>) (long dash), half (intermediate dash) and a tenth (short dash).  
 598 The error due to an uncertainty in Rpred of ±0.05 is shown and the 2003-2018 average Rpred value  
 599 (0.597±0.05) was used. The arrows indicate the change in Predicted C as Cext is reduced and Net  
 600 Cint\_2 reduces to one tenth of the current value over 21 yr.

601

602 The model indicates that the internal load needs to be reduced considerably, approaching one  
 603 tenth of the current value, to be able to achieve either G/M boundary concentration. The  
 604 main uncertainty is how the internal load reduces as Cext/Obs C (or net phosphorus  
 605 sedimentation) decreases as a result of reductions in Cext. Katsev et al. (2006) found  
 606 unstable predictions of the long-term efflux of phosphorus from sediment in their diagenetic  
 607 model when return of phosphorus from the water was included. More understanding of this  
 608 link is desirable, but we can note that there is some consistency in the time for the internal  
 609 load to reduce in the 21 lakes in Fig. 7, even though there is great variety in their  
 610 eutrophication history, size, depth, how much and how quickly the external load was reduced,  
 611 etc.

612

613 **4.2.2.** *Other lakes*

614

615 The  $t_{90}$  times for the lakes given in Table 2 were selected using the two criteria established in  
616 Section 4.2. Eight of the lakes have internal loads (Beg, Cam, Clea Lakes, Gullion, Island  
617 Reavy, Mourne, Portmore, Stoneyford) and without direct information on the rate of  
618 reduction of internal load in them, there is uncertainty in their response time. At present, this  
619 is given as 8 to 20 years, except for Lough Island Reavy. That lake has a small internal load  
620 of  $8.3 \text{ mg P m}^{-3}$  and a long  $\tau_w$  value (3.82 yr), so the  $2.303\tau_w$  criterion was applied. The  
621 remaining lakes should respond according to the  $2.303\tau_w$  values.

622

623 **4.3.** *Recovery of Lough Neagh*

624

625 The modelled recovery of Lough Neagh (Fig. 8) indicates the response of the lake  
626 concentration to reductions in the inflow concentration and net internal load. In addition to  
627 the uncertainty about how the internal load responds to reduction in inflow/lake  
628 concentration, there are two further considerations for the recovery of the lake; change in the  
629 lake phosphorus cycle from 2003 and the influence of phosphorus and nitrogen on  
630 chlorophyll *a*.

631

632 **4.3.1** *Change in the lake phosphorus cycle*

633

634 The reduction in  $R_{meas}$  from 0.330 during 1995-2002 to 0.191 in 2003-2018 (Section 3.2)  
635 represents a considerable change in the lake phosphorus cycle, and the evidence suggests that  
636 it was a consequence of the earlier depletion of total oxidized nitrogen (TON, nitrate/nitrite)

637 in the water column (McElarney et al. 2021). Once the sequence of mineralization of organic  
638 matter in the sediment moved from TON, after its depletion in the water column, to Mn(IV),  
639 then Fe(III), etc. (Froelich et al. 1979; Wersin et al. 1991), the release of phosphorus sorbed  
640 to iron oxyhydroxide from the sediment began earlier. There is no direct evidence for this  
641 sequence in the Lough Neagh sediment but there is for other lakes (Wersin et al. 1991; Park  
642 & Jaffe 1999; Hemond & Lin 2010; Smith et al. 2011). Further, movement of high soluble  
643 reactive phosphorus concentrations towards the sediment surface has been observed in Lake  
644 Okeechobee (Moore et al. 1998) and increases of soluble reactive phosphorus concentration  
645 along with soluble iron and especially manganese have been observed in the upper sediment  
646 layers of Lake Vallengtasjon (Lofgren & Bostrom 1989).

647

648 There is, however, indirect evidence to support this release mechanism in Lough Neagh.  
649 **First**, the redox potential at four sediment depths, from March to October, 1979, shows a  
650 sudden decrease in mid-July even at 10 mm, indicating that release of phosphorus was likely  
651 (Rippey & Jewson 1982); Gibson et al. (2001) measured a release of 80 tonnes P in July  
652 1979. The results for 1983 show similar behaviour, with a decrease in July, even temporarily  
653 at the sediment surface (Section 9 Supplementary materials); the release was 140 tonnes P in  
654 July 1983 (Gibson et al. 2001). **Second**, the seasonal cycle of iron, manganese and soluble  
655 reactive phosphorus in the water column of Lough Neagh, 1983 (Section 7 Supplementary  
656 materials), is similar to that described by Mayer et al. (1982) in the hypolimnion of  
657 hypereutrophic Lake Seabasticook. As mineralization proceeds to produce first soluble Mn  
658 and then Fe, the concentration of soluble reactive phosphorus increases, more closely  
659 associated with Mn than Fe. The increase of Fe in Lough Neagh is not as great as that  
660 observed in Seabasticook, probably as there is still a thin oxic layer in the sediment, compared  
661 to full anoxia in the hypolimnion of Seabasticook.

662

#### 663 4.3.2 Influence of phosphorus and nitrogen on chlorophyll *a*

664

665 Whether achieving the target inflow and lake phosphorus concentrations alone will lead to a  
666 desired chlorophyll *a* concentration needs consideration of the relative influence of nitrogen

667 and phosphorus on biological productivity (Lewis Jr et al. 2011; Schindler et al. 2016; Paerl

668 et al. 2016). The chlorophyll concentration has been reducing Lough Neagh since the mid-

669 1990s, as a result of a lower nitrogen input load and lake concentration (McElarney et al.

670 2021; Elliott et al. 2015). This is a result of catchment based nutrient management measures

671 under the Northern Ireland Nitrates Action Plan; this outlines seasonal and mass controls on

672 fertilisers in addition to restrictions on land types that can receive fertilisers and failure to

673 comply with such measures can result in financial penalties. If the trend continues, further

674 decreases in chlorophyll would be expected.

675

676 Action to reduce nitrogen inputs has been successful, whereas it is more difficult to reduce

677 phosphorus inputs from the large catchment and there is also the influence of the internal

678 phosphorus load on the lake concentration. Whether further action to reduce nitrogen or

679 phosphorus inputs or both depends on the reduction in chlorophyll expected and the relative

680 ease of reducing the inputs. A reduction in lake phosphorus concentration is, however,

681 needed to reduce the internal load.

682

683 Some indication of the response of chlorophyll to combinations of lake total phosphorus and

684 TON concentrations can be provided by using chlorophyll-nutrient models (Table 4). As no

685 Correction Factors are available for the models, a typical value of 1.15 was used, or the

686 untransformed values could be considered geometric means. Table 4 was generated using the

687 following values: observed 2003-2018 mean total phosphorus concentration of 135 mg P m<sup>-3</sup>,  
 688 72 in the absence of an internal load, 44 and 24 the target values for a colored and clear water  
 689 lake; observed 2003-2018 mean TON concentration of 294 mg N m<sup>-3</sup> and 100 which could be  
 690 produced if the current reductions continue. The observed 2003-2018 mean chlorophyll *a*  
 691 concentration is 43.4 mg m<sup>-3</sup>.

692

693 Table 4. The chlorophyll *a* concentration in Lough Neagh predicted by four models, for combinations  
 694 of total phosphorus and total oxidized nitrogen values. Model 1 is Prairie et al. (1989) Eqn 3, based  
 695 on N and P; Model 2 Prairie et al. (1989) Table 2, based on P and N/P ratio; Model 3 Phillips et al.  
 696 (2008) Eqn. 3, based on N and P; Model 4 Phillips et al. (2008) Eqn. 6, based on P for high alkalinity  
 697 shallow lakes. The predicted values could be considered geometric means and the values in  
 698 brackets are adjusted for bias. The TP/TON (mass) ratio is given.

TP, mg P m <sup>-3</sup>	TON, mg N m <sup>-3</sup>	TP/TON (mass)	Model 1	Model 2	Model 3	Model 4
135	294	2.18	9.1 (10)	25.0 (29)	28.8 (33)	34.9 (40)
72	294	4.08	6.5 (8)	16.7 (19)	17.5 (20)	20.2 (23)
44	294	6.68	5.1 (6)	12.2 (14)	11.8 (14)	13.2 (15)
24	294	12.3	3.7 (4)	6.9 (8)	7.3 (8)	7.8 (9)
135	100	0.74	3.7 (4)	25.0 (29)	20.3 (23)	34.9 (40)
100	100	1.00	3.1 (4)	20.6 (24)	16.0 (18)	26.9 (31)
80	100	1.25	2.8 (3)	17.9 (21)	13.4 (15)	22.2 (26)

699

700 Model 1 does not reproduce the chlorophyll concentration, the other models are better  
 701 predictors, with values of 29-40 mg m<sup>-3</sup>, compared to the observed 43.4, and Model 4 the  
 702 closest. A decrease of chlorophyll to 14-15 and 8-9 mg m<sup>-3</sup> as phosphorus is reduced to 44  
 703 and 24 mg P m<sup>-3</sup> is predicted. Reducing TON to 100 mg N m<sup>-3</sup> is predicted to not affect the  
 704 chlorophyll at the current total phosphorus concentration and only reduce it a little even if it  
 705 decreases (18-31 and 15-26 mg m<sup>-3</sup> for total phosphorus of 100 and 80 mg P m<sup>-3</sup>). For all of  
 706 the nitrogen and phosphorus combinations, except total phosphorus of 24 mg P m<sup>-3</sup>, the N/P  
 707 (mass) ratio still indicates nitrogen deficiency, based on phosphorus deficiency above 23 and  
 708 nitrogen deficiency below 9 (Guildford & Hecky 2000). Even at the lowest phosphorus

709 concentration, there is still no phosphorus deficiency, with both nutrients influencing the  
710 chlorophyll concentration. Further assessment is required to determine the influence of  
711 nitrogen and phosphorus on chlorophyll and other biological properties and the catchment  
712 actions needed to reduce the nutrient input loads.

713

714 The predictions of these empirical models are most accurate for lakes with typical relative  
715 concentrations of nitrogen and phosphorus. Lough Neagh, however, has an unusual  
716 relationship, as the 2003-2018 averages, 135 mg P m<sup>-3</sup> for total phosphorus, 294 mg N m<sup>-3</sup> for  
717 total oxidized nitrogen and 43.4 mg m<sup>-3</sup> for chlorophyll *a*, place it well outside the scatter plot  
718 of European lakes (Poikane et al. 2022).

719

#### 720 **Declaration of interests**

721

722 The authors declare that they have no known competing financial interests or personal  
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724

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726

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737

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